Descriptive Notes

Physiography and drainage
The main physiographic feature in this map area is the junction of the Takhini and Yukon river valleys. North of the Takhini River is the Yukon Plateau and more specifically, the southern extent of the Miners Range. South of the Takhini River is the Boundary Ranges of the Coast Mountains.

The Yukon River captures the Takhini River and flows northward into Lake Laberge. A prominent delta of the Yukon River is built out into Lake Laberge. Another important drainage in the map area is Flat Creek which flows southward into the Takhini River.

The McConnell Glaciation in the Whitehorse area
During the Late Wisconsinan McConnell Glaciation (~20 000 years ago), the Whitehorse map area (NTS 105D) was glaciated by ice lobes originating in the Coast Mountains and the Cassiar Mountains of southern Yukon. Initial ice accumulations in the map area probably began in the higher regions of the Coast Mountains. It was not until localized ice caps had formed that the more distal Cassiar Lobe advanced into the map area from the southeast through Marsh Lake valley. The convergence of the two lobes at glacial maximum occurred over the Coast Mountains west of the city of Whitehorse. At the height of the last glaciation, movement of ice over this area was to the northwest and was unobstructed by topography. Adjacent to the Alaska Highway near the hamlet of Ibex Valley, evidence of ice-flow directions from this period of the glaciation are recorded on polished bedrock surfaces as glacial striae.

The pattern of deglaciation is highlighted by periods of differential retreat and fluctuating ice fronts. A re-advance of the Cassiar Lobe occurred into this area and had a significant influence on sediment deposition on the landscape. In particular, systems of pro-glacial lakes developed marginally to the Cassiar Lobe in the Takhini River and Lake Laberge valleys as it retreated. The glacial lake in the Takhini River valley, also known as Glacial Lake Champagne, during its latest stage drained northward via the headwaters of the Nordenskold River. This outlet maintained a maximum water level of 746 m a.s.l. or about 100 m (328ft.) depth. Glacial Lake Laberge developed behind a sediment dam at the far north end of the present lake (this natural dam still controls the water level in the lake). The outlet through the sediment dam has since dropped due to erosion but at its maximum level the depth of Glacial Lake Laberge was 72 m (239 ft.) above the current shoreline. The water depth over the city of Whitehorse was likely <5 m deep.

Landforms
Takhini River valley
As the glacier front retreated to the east out of the Takhini River valley, it periodically paused and deposited large end moraines (Figure 1; see the moraine symbol on the map). The curvature of these landforms reflects the shape of the ice margin at the time of deposition. These moraines are a good aggregate resource in a valley bottom that is otherwise devoid of coarse sediment (see station JB-03-002 on map).

The dominant landforms in the Takhini River valley are the plains that once formed the bottom of Glacial Lake Champagne. The Takhini River has since down-cut into these silty deposits (Figure 2; see "LG" on the map) and exposed the depositional layering from within the former lake. Side tributaries entering the Takhini River from the south also cut impressive canyons into the glacial lake bed sediments.

Scout Lake valley
The glacier that filled the Takhini River valley spilled into the Scout Lake valley. As this glacier retreated large meltwater channels were carved into the valley sides (Figure 3; see meltwater channel symbols). Many of these channels are very large and attest to the quantity and force of water draining off the melting glaciers.

Lake Laberge and Yukon River valley
Like the Takhini River valley the dominant landforms in the Yukon River valley are glacial lake plains. On the west side of the Yukon River these plains are visible from the North Klondike Highway. The reduction in lake levels in this valley has been ongoing since their initial high stand following glacial retreat to the present day level of Lake Laberge. This is due to the continual erosion of the sediment dam at the north end of Lake Laberge. Evidence of this progressive lake level reduction is visible along the North Klondike Highway between Takhini Hot Springs Road and Shallow Bay Road. The highway traverses over terrace edges which represent former shorelines of Lake Laberge (see strandline symbol). As you drive north, the shorelines become lower in elevation until you reach the modern shoreline at Shallow Bay. Shorelines are also visible immediately north of the Horse Creek road near its junction with the North Klondike Highway. Other evidence of former shorelines is preserved as beach ridge dunes on the east side of the Yukon River near Swan Lake and the Livingstone Trail. These beaches formed on a south-facing shoreline that was part of a former bay of Lake Laberge, similar to Shallow Bay today. This shoreline has an age of 10000 +/- 1000 years old. A sand spit that developed from river currents entering this former bay is well preserved in Swan Lake and forms a point out into the lake (Figure 4).

Following glacial retreat, the reduction in lake levels in the Yukon River valley exposed the former lake bottom surface. During this time the winds coming off the retreating glaciers to the south were strong and began reworking the lake bottom sediments into parabolic dunes (see Eolian materials on the map). The last period of eolian activity (active dunes) in the Whitehorse dune field was approximately 9100 +/- 1000 years ago. At this time the wind strength probably waned and vegetation stabilized the land surface. While the largest dunes are located east of the Yukon River near the Livingstone Trail there are also smaller dunes visible in North Klondike Highway road cuts, immediately north of the Takhini Road junction.

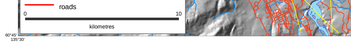
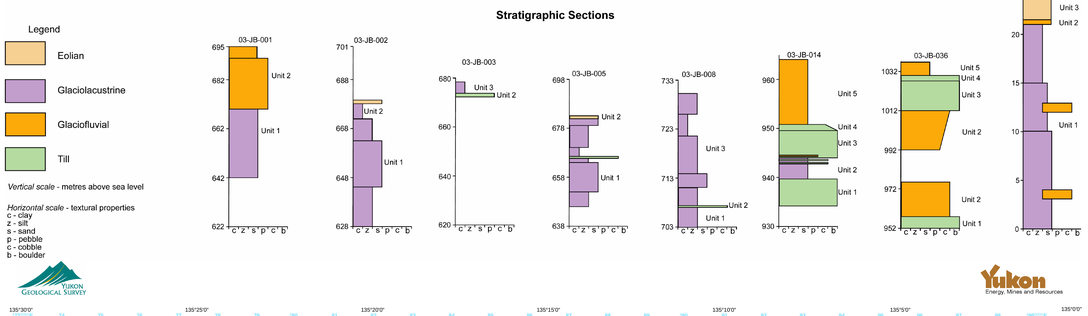


Figure 1. A view south across the Takhini River valley towards Haeckel Hill. A recessional moraine and its well-drained surface is visible in the foreground. The moraine ridge is visible across the valley. The terrace surface to the left of the moraine is made up of glaciolacustrine plains.

Figure 2. An aerial view to the west up the Takhini River valley. The glaciolacustrine plains are exposed by down-cutting of the Takhini River. The fine-grained fluvial deposits that make up the Takhini River floodplain have been developed for agricultural purposes.

Figure 3. An aerial view to the northeast from the Scout Lake valley. Scout Lake is visible in the background left. The deeply incised channels are remnant of ice-marginal glacial meltwater erosion. These channels were cut when a tongue of ice from the Takhini River valley spilled into the Scout Lake valley during glacial recession.

Figure 4. An aerial view looking down on Swan Lake. The Yukon River is visible in the background. The point of land extending out into Swan Lake is a relict sand spit. It was active when the water level in Lake Laberge was higher and the southern shoreline of the lake was positioned in this area. Swan Lake would have been a bay that was connected to Lake Laberge. The sand spit is remnant of a water current that was depositing sand into this former bay.

Stratigraphic Sections